

Preface: Rift tectonics and syngenetic sedimentation – the Cenozoic Lower Rhine Basin and related structures

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The collection of extended and updated abstracts presented here resulted from a two-day workshop of the Collaborative Research Centre 'Interactions between and Modelling of Continental Geo-Systems' (SFB 350) held on 19th and 20th November, 1999 at the Rheinische Friedrich-Wilhelms-Universität Bonn. This workshop was entitled 'Rift Tectonics and Syngenetic Sedimentation – the Cenozoic Lower Rhine Graben and Related Structures'. Its central topic was devoted to various approaches of understanding the evolution of a young graben structure, with the aim to trace the interaction of tectonics, sedimentation, facies development, burial history, and coherent structural behaviour by integration of geodetic, seismological, gravimetric, sedimentological, paleontological, and climatological data. Analysis and modelling of sedimentary basins, applied to an area of intense intercalation of marine and terrestrial sedimentation, only works, if there are sufficient and consistent stratigraphical data to precise the age control of the deposits.

The papers have different scopes, but overall they show that the Roer Valley Graben (= Rur Graben) and the Lower Rhine Basin, parts of a major Cenozoic rift system in the Netherlands and Germany, seem to be – in many details – an astonishingly little-known example of a rift-related graben structure forming a sedimentary basin in the very centre of Europe.

The Dutch-German rift system (Fig. 1) transects the Rhenish Shield ('Rheinischer Schild' sensu Cloos, 1939), which comprises the Rhenish Massif ('Rheini-

schen Schiefergebirge') in the north and the ancient crystalline massifs of the Black Forest ('Schwarzwald') in the south of Germany and the Vosges ('Vogesen') in neighbouring France. The Upper Rhine Graben bifurcates at the southern margin of the Rhenish Massif NW-ward into the Lower Rhine Basin and the Roer Valley Graben and NE-ward into the Hessen depression, forming a triple junction (Ziegler, 1994). The NW-SE trending faults of the Lower Rhine Basin developed from Early Oligocene times onward, overprinting the pre-existing Paleozoic and Mesozoic structural framework (Vinken, 1991). The Lower Rhine Basin deepens toward the Roer Valley Graben in the Netherlands (Verbeek et al., 2002), passing with an offset into the West Netherlands Basin (Cloetingh et al., 2001) and offshore into the Broad Fourteens Basin (Wong et al., 2001).

The southern part of the Dutch-German rift system cuts into the north-western headlands of the Rhenish Massif, forming the 100 km long and 50 km wide lowland area, the Lower Rhine Embayment in Germany ('Niederrheinische Bucht'). The rift system consists of a mosaic of north-eastward tilted half-grabens and horst blocks separated by NW-SE-trending and SW-ward dipping synthetic normal and conjugate NE-ward dipping antithetic faults. Within these grabens, the maximum thickness of the Tertiary and Quaternary siliciclastic basinfill is attained at the respective north-eastern boundary faults, e.g. for the Roer Valley Graben up to 2000 m along the SW side of the Peel Fault and for the Lower Rhine Basin up to

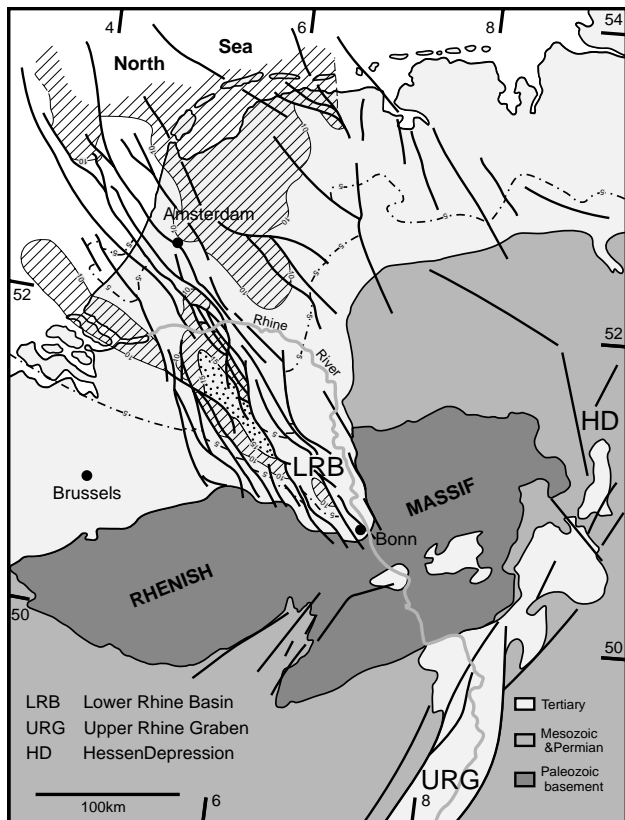


Fig. 1. Structural map verifying the base of the Tertiary in Belgium, the Netherlands, and northern Germany, following Vinken (1991), Zagwijn et al. (1985), and Hager & Prüfert (1988). The Dutch-German rift system is narrow in the south-east and widens toward the north-west, forming a fan of diverging faults. The overall structure of the rift system might be formed by extension due to sinistral transtensional shear. The Roer Valley Graben and the Lower Rhine Basin could be separated by conjugate shear to form the two individual Dutch and German sub-basins.

1300 m along the SW side of the Erft Fault. Rich lignite seams (forming the Main Lignite Seam with a maximum thickness of about 100 m in the SE part of the Erft Block) have intensely been mined for decades by the local lignite coal industry in open-cast mines in the Lower Rhine Basin. The detailed spatial information on strata and fault geometry available from the mining survey provided the basis for three-dimensional models of backward reconstructions of sedimentation and compaction (Jentzsch & Siehl, 2002). Moreover, also the design of a kinematic model of a typical faulted domain with the objective to upscale the structural development to a larger area within the rift system could be made (Thomsen & Siehl, 2002). To facilitate the design and implementation of 3D/4D geological applications that can hardly be modelled within standard GIS, a component software called *GeoToolKit* was developed (Breunig et al., 2002).

During the Cenozoic, the embayment area subsided tectonically prior to lignite mining at rates of about 2 mm/a along the north-eastern margins of the

individual main graben blocks (Quitow & Vahlen-sieck, 1955). Maximum subsidence, including an anthropogenic component due to mine water drainage, amounts to the present rate of at least 30 mm/a, in some places even more. A horizontal extension in east-west direction of up to 2 mm/a is inferred from GPS measurements (Campbell et al., 2002). Regional distribution of recent seismic activity points to a close association with the main normal fault systems, where strike-slip components have been under debate (Ahorner, 1975, 1983, 1997; Klostermann, 1983; Plein et al., 1982; Schreiber & Rotsch, 1998). Strike-slip components caused by a sinistral transtensional opening of the entire Dutch-German rift system may indeed be indicated by the eastward tilt of its strata and the eastward offset of its southern part, the Lower Rhine Basin.

From the Paleogene onward, the North Sea transgressed onto the lowlands of present-day Belgium, the Netherlands, and Germany. During the initial rift subsidence in Oligocene times – following the rift structures from the north (Kockel, 2002) – the sea overlapped the ‘Schiefergebirge’ and reworked the debris which was shed by fluvial input from its heights. The ingress of the sea is well documented by coastal marine environments which were deposited under tidal conditions. After the initial rifting during the Early Oligocene, younger basin-fill units apparently prograded northward due to a structural shift of the basin depocentre. The interaction of the Cenozoic North Sea with the coastal plains of the Lower Rhine Basin resulted in well-documented cyclic sedimentation. In the Early Miocene, a coastal plain was formed, that was covered by swamps and forests. Sedimentological analyses from wells and open-cast mines (Klett et al., 2002) and a large amount of biostratigraphical data from within and outside the basin (Moers, 2002; Heumann & Litt, 2002), yielded information about its sedimentary development and the timing of structural evolution.

When the North Sea retreated again in the Late Miocene and Pliocene, fluvial systems followed the seaward progradation of the coast, and the early precursors of the recent rivers Meuse, Rhine, Sieg, and Wupper frequently cut across the subsiding basin (Van Balen et al., 2002). A gradual decrease in temperature was documented by changes in the floral composition of the plant communities (Utescher et al., 2002). Finally, the Rhenish Massif was uplifted in the Pleistocene to reach its present elevation, and the Rhine River deposited a thick and coarse-grained fluvial outwash-fan towards the North (Boenigk, 2002). In the Rhenish Massif, a staircase of terraces developed by fluvial incision and aggradation as a result of

the complex interaction of tectonics, sea level, and climatic changes (Meyer & Stets, 2002).

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